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A new fractal algorithm to model discrete sequences*

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Employing the properties of the affine mappings, a very novel fractal model scheme based on the iterative function system is proposed. We obtain the vertical scaling factors by a set of the middle points in each affine transform, solving the difficulty in determining the vertical scaling factors, one of the most difficult challenges faced by the fractal interpolation. The proposed method is carried out by interpolating the known attractor and the real discrete sequences from seismic data. The results show that a great accuracy in reconstruction of the known attractor and seismic profile is found, leading to a significant improvement over other fractal interpolation schemes.

Keywords: fractal interpolation, the vertical scaling factors, iterative function system, seismic data

PACC: 0555

1. Introduction

There are many objects in nature which are so complicated and irregular that they cannot be modeled properly with the classical geometry. When classical geometry fails to serve as a tool to analyse the complexity of such object, fractal geometry begins. The concept of fractal geometry was first introduced as an extension of classical geometry by Mandelbrot, which can be used to make accurate models of physical structures from ferns to galaxies.^[1]

A fractal object is self similar in that subsections of the object are similar, in some sense, to the whole object. No matter how small a subdivision is taken, the subsection contains details no less than the whole.^[2] The fractal dimension, thus, is introduced as a measure of the scaling property of the features. Therefore, any irregular shaped body, including signals and discrete sequences, whose Hausdorff–Besicovitch dimension strictly exceeds the topological dimension, can be measured by the fractal dimension:^[3,4] many natural shapes such as coastlines, mountains and clouds are easily described by fractal models.

Many signals or discrete sequences, such as stock price, seismic data, etc, are scale invariance, thus, the best way to model such signals or discrete sequences is to use a fractal model.^[5] There are two popular ways

in which a fractal object can be constructed. One is in terms of fractional Brownian motion, which is statistically self similar. The other involves the iterated function system (IFS) theory,^[6] which is a more general approach than fractional Brownian motion.^[5]

The IFS theory has received a great deal of attention.^[7–9] Since IFSs are capable of generating complicated and varied functions even if maps, as few as two ($m = 2$), are involved, we can find their applications in many fields, such as signal processing, computer graphics, metallurgy, geology, chemistry, medical sciences and other areas.^[10–17] However, one of the challenges of using IFS for modeling the discrete sequences is the difficulty in determining the vertical scaling factors. In Ref. [17], an algorithm of the computation of the vertical scaling factor was introduced based on the local information. This method, however, needs to employ a great many of known points to achieve a desired data model accuracy because it actually is a linear interpolation. The rest of the present paper is organised as follows. In Section 2, after analysing the IFS based on the affine transformation, we propose a very novel and simple method to compute the vertical scaling factors, in which used are only the uniformly distributed known points in the discrete sequences. The underlying fundamental behind the algorithm is mentioned in Section 3. The proposed method is used to model two discrete sequences

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The vertical length between (x_h, y_0^1) and (x_h, y_0^0) is Δy_0 , however it becomes Δy_n by the affine mapping ω_n . According to fundamental of the affine mapping, the relationship can be constructed by

$$\Delta y_n = d_n \times \Delta y_0. \quad (9)$$

Therefore, the vertical scaling factor corresponding to ω_n can be defined as

$$d_n = \frac{\Delta y_0}{\Delta y_n}, \quad n \in [1, 2, \dots, N]. \quad (10)$$

4. Simulations

To testify the reliability and the validity of the proposed fractal modeling method, we give two numerical examples: one is to reconstruct a known attractor and the other is to try to interpolate the discrete sequence-the recorded field seismic data.

4.1. Reconstruction of known attractor

The graph of the interpolation function is $\mathbb{G} = \{(x, 2x - x^2) : x \in [0, 2]\}$. According to the proposed method in the paper, we can claim that $\mathbb{G}$ is the attractor of the hyperbolic IFS $\{\mathbb{R}^2; \omega_1, \omega_2\}$ if we choose the data set of $\{(0, 0), (1, 1), (2, 0)\}$ as the interpolation points. Therefore, the two corresponding affine mappings can be expressed as

$$\begin{aligned} \omega_1 \begin{bmatrix} x \\ y \end{bmatrix} &= \begin{bmatrix} 0.5 & 0 \\ 0.5 & 0.25 \end{bmatrix} \begin{bmatrix} x \\ y \end{bmatrix}, \\ \omega_2 \begin{bmatrix} x \\ y \end{bmatrix} &= \begin{bmatrix} 0.5 & 0 \\ -0.5 & 0.25 \end{bmatrix} \begin{bmatrix} x \\ y \end{bmatrix} + \begin{bmatrix} 1 \\ 1 \end{bmatrix}. \end{aligned} \quad (11)$$

As x varies over $[0, 2]$, the term on the right-hand side of the first equation in expression (11) yields a part of the graph of $f(x)$ lying in self-affine region $[0, 1]$, while the term on the right-hand side of the second equation yields a part of the graph of $f(x)$ lying in interval $[1, 2]$. Hence $\mathbb{G} = \omega_1(\mathbb{G}) \cup \omega_2(\mathbb{G})$. Since $\mathbb{G} \in \mathbb{H}(\mathbb{R}^2)$ we conclude that it is the attractor of the IFS by noticing that the IFS is just-touching. The reconstructed known attractor $\mathbb{G}$ by the proposed method is shown in Fig. 2. For comparison, the same attractor reconstructed by the method in Ref. [17] is shown in Fig. 3. From the interpolation errors for these two methods shown in Fig. 4, we can see that the error for the proposed method is much less than that for the compared one.

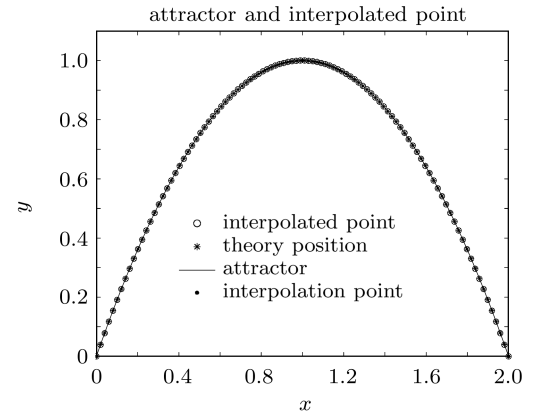


Fig. 2. Reconstructed known attractor by proposed method.

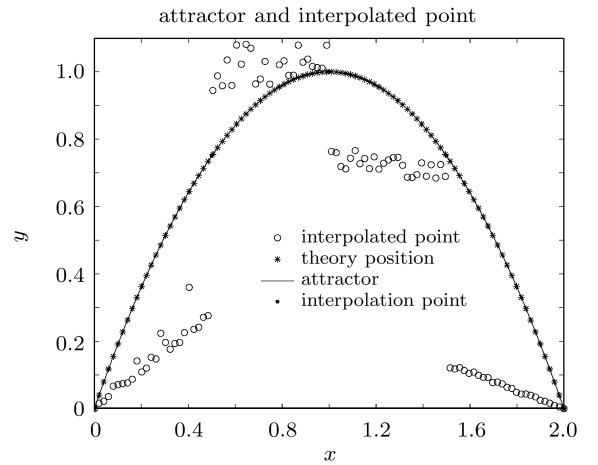


Fig. 3. Reconstructed known attractor by compared method.

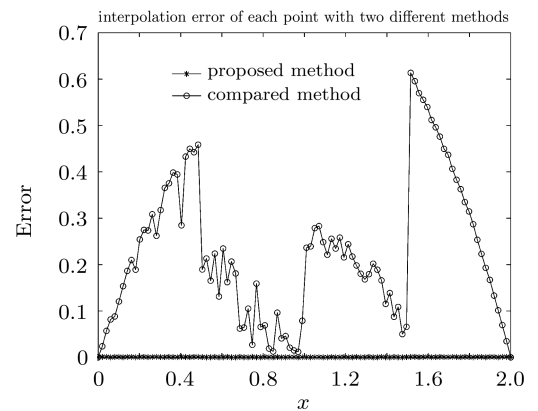


Fig. 4. Reconstruction errors induced by proposed and compared methods.

4.2. Reconstruction of the real discrete sequence

In the second simulation, the recorded field seismic data are considered as the graph of the IFS, and

we try to obtain this graph by the proposed method with the following procedure.

(1) Select a set of points $(x_n, y_n), n \in [1, 2, \dots, N]$ as the interpolation point grouped into the set $\mathbb{G}$ that has been used in the Eq. (5): $\mathbb{G} = \{(x_n, y_n), n \in [1, 2, \dots, N]\}$. The points selected as the interpolation points can be distributed uniformly or not in the sequence. In our test, the interpolation points are uniformly distributed.

(2) Compute the affine mappings $\omega_n, n \in [1, 2, \dots, N]$ according to Eqs. (8) and (10).

(3) Substitute the set $\mathbb{G} = \{(x_n, y_n), n \in [1, 2, \dots, N]\}$ and the affine mappings $\omega_n, n \in [1, 2, \dots, N]$ into Eq. (5). A new set of interpolation points will be obtained and the procedure can be continued until the interpolation points are dense enough to obtain the expected interpolation resolution. Note that once the affine mappings are obtained from the second step, they will not change with the interpolation points obtained from later circulations.

The reconstruction result can be seen in Fig. 5. In this figure, the reconstructed traces numbers are 10, 20, 40, 65, 80, and they match the whole profile quite well. Figure 6 shows the original and the reconstructed waveforms obtained by the two methods. In general, these two methods can recover the main features of traces. However, if the traces are magnified, the difference between the proposed and the compared method becomes quite obvious, which can be seen in Fig. 7. The trace recovered by the compared method is not very smooth, while the one by the proposed method is as smooth as the original one. This ability is quite important in the modern seismic processing because some other methods need dense and regular sampling traces, such as surface-related multiple elimination and wave-equation migration. The improvement by the proposed method in terms of interpolation error can also be seen in Fig. 8. The error for each point, induced by the proposed method, is much smaller than that induced by the compared method, while the interpolation errors for each trace are 2 or 3 times lower than that induced by the method in Ref. [17], see Tabel 1.

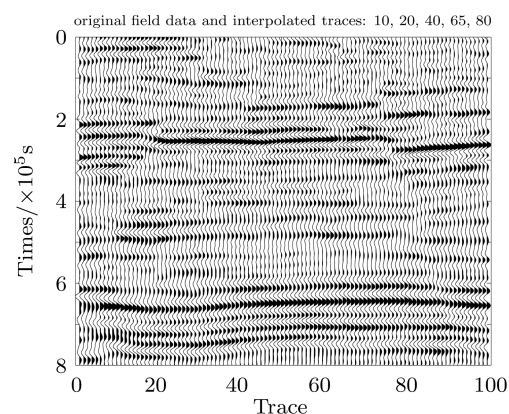


Fig. 5. Reconstructed seismic data profiles obtained by proposed method. Reconstructed trace numbers are 10, 20, 40, 65 and 80.

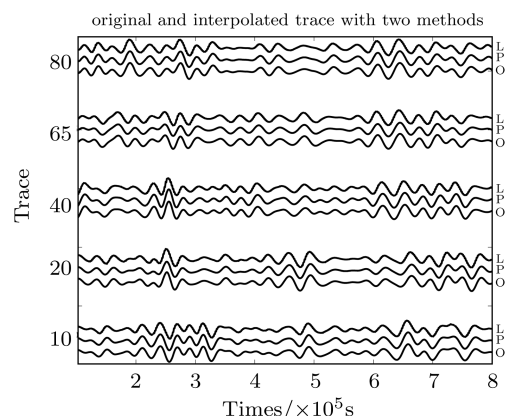


Fig. 6. Reconstructed traces obtained by proposed and compared methods. The corresponding trace numbers are also 10, 20, 40, 65 and 80. L means reconstructed sequence obtained by compared method, P represents that by proposed method, and O refers to original sequence.

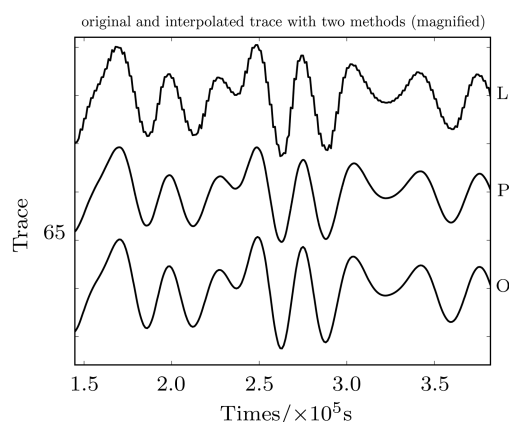
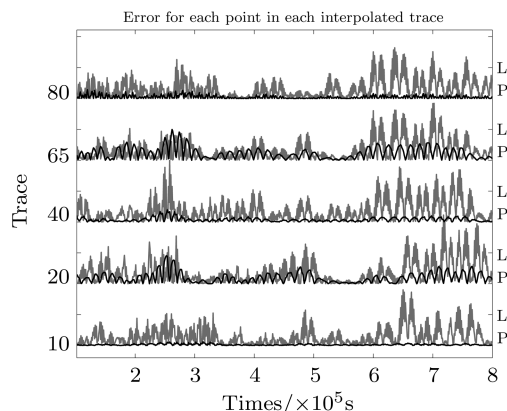


Fig. 7. Magnified part of reconstructed traces obtained by proposed and compared methods. The corresponding trace numbers are also 10, 20, 40, 65 and 80. L means reconstructed sequence obtained by compared method, P represents that by proposed method, and O refers to original sequence.

Table 1. Errors for each reconstructed trace by two methods.

trace number	10	20	40	65	80
the proposed method	24.3636	179.7889	70.6344	209.4774	44.6208
the compared method	292.6753	382.815	331.5845	296.1738	278.1165

**Fig. 8.** Reconstruction errors induced by proposed and compared methods for reconstructed trace numbers 10, 20, 40, 65 and 80. L means reconstructed sequence obtained by compared method, P represents that by proposed method, and O refers to original sequence.

5. Conclusions

In this paper we derive an efficient algorithm of performing IFS interpolation and parameter estimation, say the vertical scaling factor in each affine mapping, for a given self-affine signal. We emphasise goodness-of-fit to the given signal. The algorithms are applied to two examples of self-affine signals. The simulation results show that the robust algorithm is strong enough to converge to the true result when the alternative method cannot do so. Real field seismic data are also used to test the algorithm, and the results show that the proposed algorithm achieves better fidelity to data than the alternative algorithm. The results also illustrate that the proposed interpolation algorithm yields more significant improvement over methods suggested in the previous literature.

References

- [1] Mandelbrot B 1982 *The Fractal Geometry of Nature* (San Francisco: W. H. Freeman and Co.) p31–60
- [2] Mazel D S and Hayes M H 1992 *IEEE Trans. Signal Proc.* **40** 1724
- [3] Unser M and Thierry Blu 2007 *IEEE Trans. Signal Proc.* **55** 1352
- [4] Unser M and Thierry Blu 2007 *IEEE Trans. Signal Proc.* **55** 1364
- [5] Zhu X, Cheng B and Titterton D M 1994 *IEE Proc-VIS Image Sign. Proc.* **141** 318
- [6] Barnsley M F 1988 *Fractals Everywhere* (New York: Academic) p205–290
- [7] Mazel D S 1994 *IEEE Trans. Sign. Proc. E* **42** 3269
- [8] Price J R and Hayes M H 1998 *IEEE Sign. Proc. Lett.* **5** 228
- [9] Craciunescu O I, Das S K, Poulson J M and Samulski T V 2001 *IEEE Trans. Biomed. Eng.* **48** 462
- [10] Wang M J and Wang X Y 2009 *Acta Phys. Sin.* **58** 1467 (in Chinese)
- [11] Yang J, Lai X M, Peng G, Bian B M and Lu J 2009 *Acta Phys. Sin.* **58** 3008 (in Chinese)
- [12] Yang X D, Ning X B, He A J and Du S D 2008 *Acta Phys. Sin.* **57** 1514 (in Chinese)
- [13] Gemperline M C and Siller T J 2002 *J. Comput. Civil Eng.* **16** 184
- [14] Nath S K and Dewangan P 2002 *Geophys. Prospect.* **50** 341
- [15] Tirosh S, Van De Ville D and Unser M 2006 *IEEE Trans. Image Proc.* **15** 2616
- [16] Bouboulis P and Dalla L 2007 *J. Math. Anal. Appl.* **336** 919
- [17] Li X F and Li X F 2008 *Chin. Phys. Lett.* **25** 1157